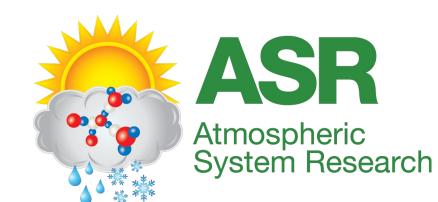
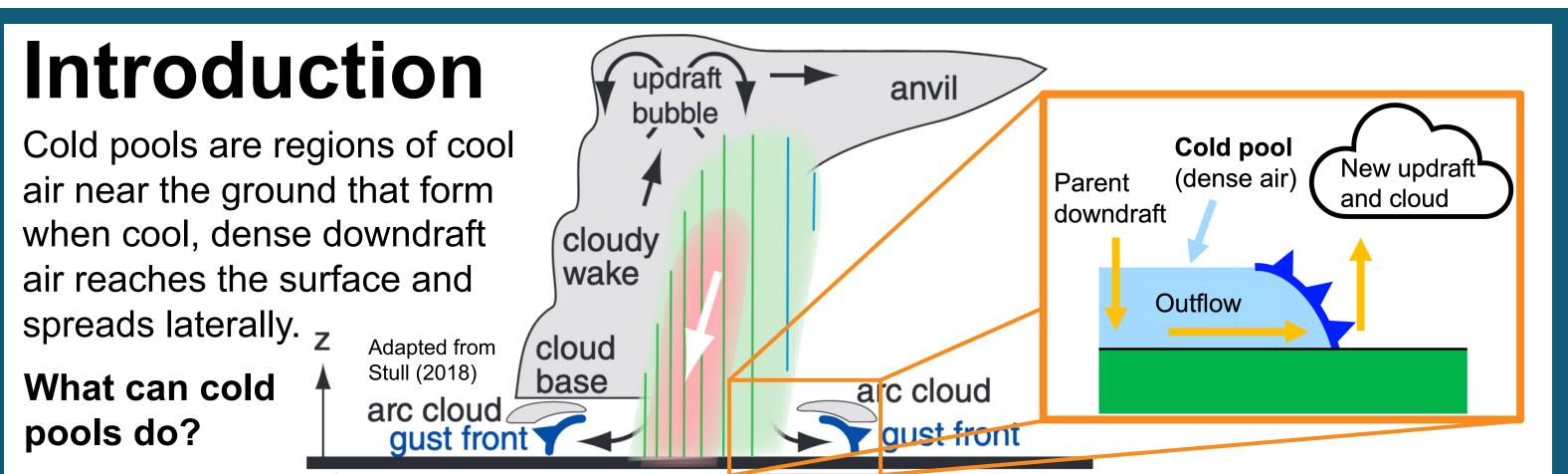
# Effects of Sea Breezes and Cold Pools on Convective Evolution during TRACER

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0600 UTC (start)





- Trigger new updrafts
- Suppress convection within their stable interiors
- Promote storm organization and longevity
- Alter the thermodynamic and dynamic characteristics of the boundary layer (cool air, gusty winds, etc.)
- Transport aerosol particles and trace gases
- Prolong the diurnal cycle of convection (old storms can trigger new storms that would not otherwise form)
- Modulate processes such as convective aggregation, tropical cyclogenesis, and tornadogenesis

### Why are cold pools cold?

Traditional explanation:

• Phase changes in the rain shaft, such as evaporation, melting, and sublimation, cause latent cooling of downdraft air.

## This study asks: **Do rainfall-land surface** interactions help to fuel cold pools?

- How do interactions between rainfall, land surfaces, and near-surface air affect cold pool evolution?
- What are the downstream effects on convective organization and interactions with the sea breeze?

## What happens to rainfall after it reaches the land surface?

- Interception by vegetation
- Shedding and stemflow (plants → soil)
- Contributes to soil moisture
- Surface water
- Runoff

#### Consequences:

- Cools and moistens the near-surface air:
  - Alters the partitioning of surface fluxes, promoting latent heat fluxes at the expense of sensible heat fluxes
  - Increased evaporation (of moister soil and vegetation-intercepted rain)
  - Increased transpiration (reduced heat stress, increased soil moisture)
- Reduces surface temperature through direct cooling (cold rainwater)
- May also cause surface warming if soil moistening lowers the soil albedo

## Mechanism-Denial Experiments

NOTE: As used here, "puddles" refers to all rainfall-land surface interactions.

1830 UTC (branching time)

Regional Atmospheric Modeling System (RAMS) simulations:

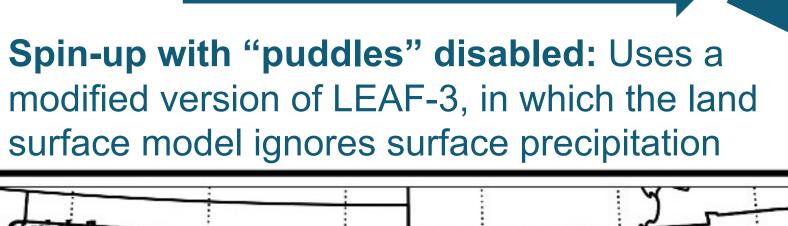
17 June 2022 case study (observed during TRACER)

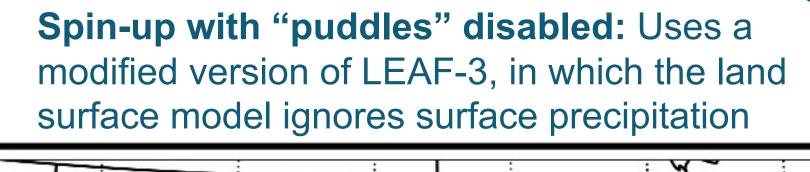
**Based on TRACER Model Intercomparison Project** (TRACER-MIP) setup by Fan, Saleeby, and others: https://arm-synergy.github.io/tracer-mip/Roadmap.html

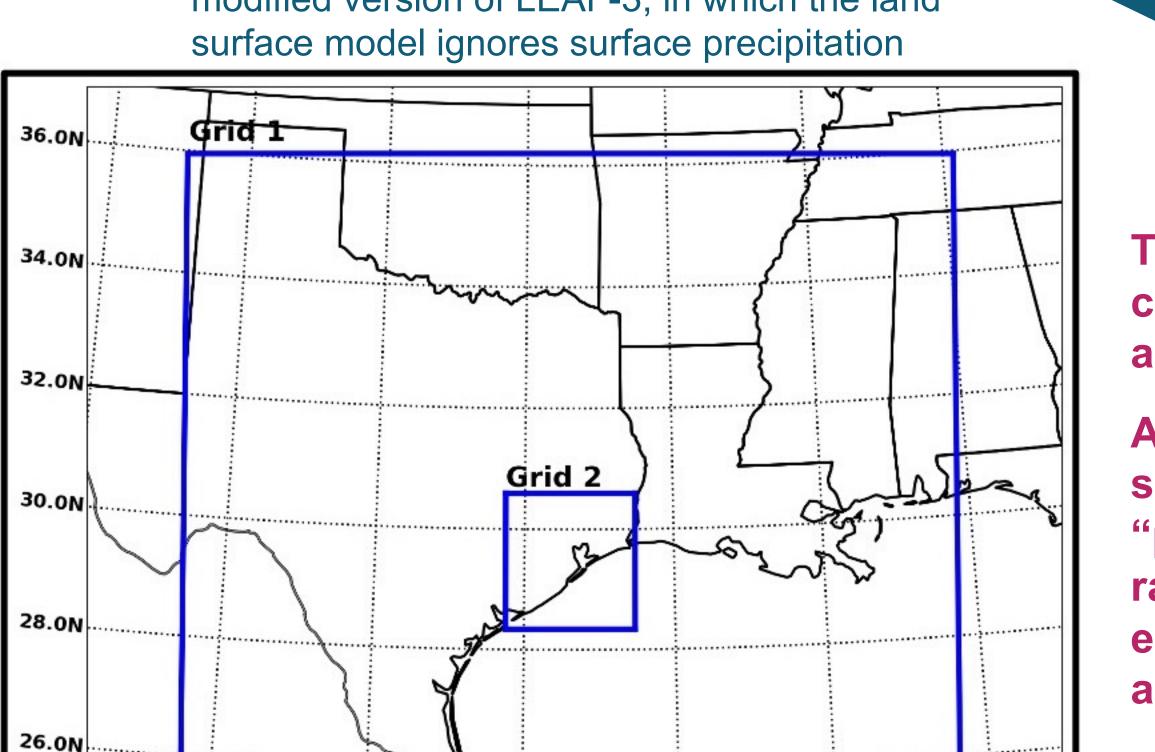
## Model grids shown at right →

22.0N

- Horizontal grid spacing: outer grid: 2 km, inner grid: 500 m
- Vertical grid: 95 stretched vertical levels (~50 m spacing near surface)
- Initial and boundary conditions: ERA5 (3-hourly)
- Surface: Land Ecosystem-Atmosphere Feedback model, version 3 (LEAF-3) with modifications
- Microphysics: RAMS two-moment bin-emulating bulk microphysics, 7 hydrometeor classes
- Radiation: Harrington (1997) two-stream scheme
- **Diffusion:** Modified Smagorinsky (1963) scheme
- Aerosol: TRACER-MIP setup; sources and sinks active; radiatively inactive







NO PUDDLES: Continue the original simulation

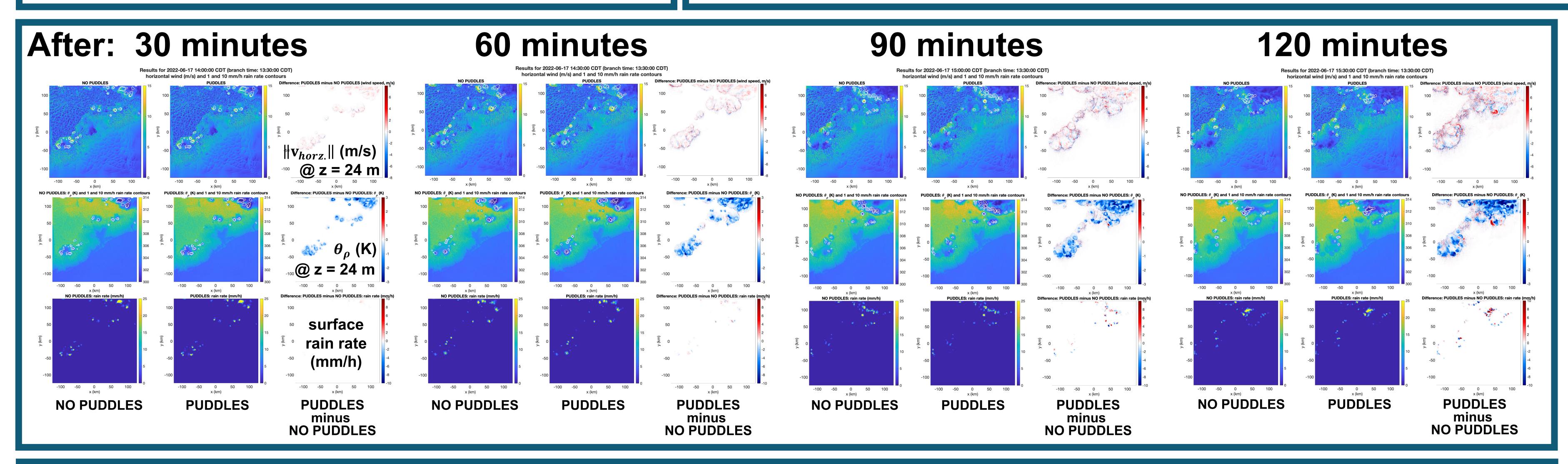
...then continue for 2 hours

**PUDDLES:** Turn on puddles (revert to the standard version of LEAF-3)

The "branching time" occurs after clouds have already begun to form and precipitate.

At the restart time, the existing simulation branches into two "parallel universes" in which rainfall-land surface interactions either remain disabled or are abruptly reenabled.

The differences that gradually emerge between these "parallel universes" (particularly within the first ~90 minutes) reveal the impact of rainfall-land surface interactions on cloud and cold pool evolution.



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- Fallen rainfall invigorates cold pools in the coastal southeast Texas setting where TRACER took place
- Conclusions: Downstream effects on convection and precipitation are observed in these simulations
  - Next steps: perform additional mechanism denial tests (rain shaft cooling, sea-breeze formation) and additional case studies



